

# Integration of geoelectrical, hydrochemical and geothermal data to identify the origin of thermal water in the foothills of the Tlemcen mountains, Northwestern Algeria

## Integrazione di dati geoelettrici, idrochimici e geotermici per identificare l'origine dell'acqua termale ai piedi dei monti Tlemcen, Algeria nord-occidentale

YOUSFI Somia<sup>a</sup> ✉, KERZABI Rachid<sup>a</sup>, MUDRY Jacques<sup>b</sup>, BENSALAH Mustapha<sup>a</sup>, ACHACHI Abdelhamid<sup>c</sup>, COLLIGNON Bernard<sup>d</sup>

<sup>a</sup>Laboratoire n°25 « PRHPM-LE-CT », Université de Tlemcen - Algérie - email ✉ : [sou.yousfi@gmail.com](mailto:sou.yousfi@gmail.com)

<sup>b</sup>Université de Franche-Comté, CNRS, Chrono-environnement, F-25000 Besançon, France.

<sup>c</sup>Agence Nationale de la Ressource Hydrique Oran- Algérie.

<sup>d</sup>Bottombillion, Avignon - France

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### Correspondence to:

Yousfi Somia ✉  
[sou.yousfi@gmail.com](mailto:sou.yousfi@gmail.com)

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### Riassunto

In precedenti studi, il gradiente geotermico medio della regione di Tlemcen è stato stimato essere basso (3,7°C/100m). I pozzi trivellati nella piana di Remchi (a Nord delle montagne di Tlemcen) rivelano la presenza di acque termali (26-48°C) e un notevole artesianismo. Per comprendere l'idrodinamica e l'origine del termalismo di queste acque sotterranee, abbiamo analizzato congiuntamente i dati geofisici e i parametri fisico-chimici; i dati geofisici ci hanno permesso di migliorare la risoluzione della rete di faglie, di stimare lo spessore dei vari letti di arenaria connessi con il serbatoio idrotermale e di delimitare l'acquifero produttivo. Le caratteristiche chimiche delle acque sono state utilizzate per un'analisi statistica che ha mostrato la presenza di due cluster di acque indipendenti derivanti da due comportamenti idrodinamici nell'acquifero. L'analisi idrotermale ha identificato la profondità raggiunta dalle acque e la relazione tra la circolazione idrica e la rete di faglie. Il confronto con gli studi precedenti ha permesso di realizzare un modello geologico schematico, che spiega l'idrodinamica, l'origine e il ruolo di questa struttura nel flusso delle acque sotterranee.

### Abstract

*In previous studies, the mean geothermal gradient of the Tlemcen region has been estimated as low (3.7°C/100m). The existing boreholes in the Remchi plain (North of the Tlemcen Mountains) reveal thermal waters (26-48 °C) and a remarkable artesianism. To understand the hydrodynamics and origin of thermalism of this groundwater, we merged geophysical and physical-chemical data. Geophysical data allowed us to improve the resolution of the fault network, to estimate the thickness of various sandstone beds connected with the thermal reservoir and to delimit the productive aquifer. Chemical characteristics of waters have been used for a statistical analysis, which displayed two independent water clusters resulting from two hydrodynamic patterns throughout the aquifer. The hydrothermal analysis identifies the depth reached by water and the relationship between water circulation and the fault network. Comparison with previous studies allowed us to build a schematic geological model, which explains hydrodynamics, origin and role of this structure in the groundwater flow.*

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## Introduction

Algerian thermomineral springs were inventoried and studied by many researchers as early as 1842. Several recent scientific works, dealing with regional thermal waters in Algeria (Issaadi 1996, 2005; Kedaid, 2006; Fekraoui, 2007; Ouali et al., 2007; Mansour et al. 2011, Bouchareb-Haouchine et al., 2012, Boudoukha et al., 2015 and Lakrouit 2022) show that the geothermal Algerian resources are of low energy type, i.e. with temperatures lower than 90°C (BRGM, 2005). In the Carbonate Mountains of Tlemcen (NW Algeria), previous hydrogeological and geothermal studies estimated the mean geothermal gradient as 3.7 °C / 100 m (Collignon, 1986). However, recent exploitation of the Remchi plain sandstone aquifer, in the northern foothills, reveals remarkable thermal waters and artesianism. In the present work, we try to explain flow processes and groundwater origin, to better define the boundaries of the geothermal reservoirs in the Remchi plain, using the joint use of geological, structural, geophysical, hydrochemical and geothermal data.

The geoelectric method is considered an effective tool for mapping groundwater resources and identifying aquifer geometry. The mapping of horizontal and vertical variation of resistivity and the use of lithology maps, allowed us to determine the fault network and to identify three zones characterised by different depths and thicknesses of sandstone layers. In addition, we used Piper-Kelly and Duvor's diagram to determine the chemical facies and eventual water groups. However, understanding the origin of groundwater requires the statistical treatment of chemical samples.

## Materials and Methods

### Study area

The study area is located in the North West of Algeria and occupies the Northern foothills of the Tlemcen Mountains (Western part of the Tell Atlas); it is limited to the North by the Sebâa Chioukh ridge, to the East by the Isser wadi, and to the West by the Tafna wadi.

Located in the middle of Tafna, the study area is covered by the watersheds of Boumessauoud, Zitoun, Sikkak and Isser wadis (Fig. 1), all originated from the Jurassic formations of the Tlemcen Mountains. It joins the Tafna wadi towards the North, crossing the Remchi plain.

The climatic regime in this area is similar to that of the semiarid Mediterranean areas of Northern Algeria (Meddi et al., 2010). The precipitations range from 160 to 1100 mm/year, as observed in ten weather stations of the middle Tafna watershed (Belarbi, 2018; Amar Belarbi & Bensliman, 2023).

### Geological and hydrogeological setting

The Tlemcen Mountains and their Northern foothills are characterised by two geological domains: the karstic Jurassic carbonate massif and the Miocene formations filling the large northern basin (Fig. 2). The first domain includes three main aquifers (Bensaoula, 2007), from the bottom to the top:

- Boumedién sandstone aquifer: 500m of moderately pored and fissured sandstones;

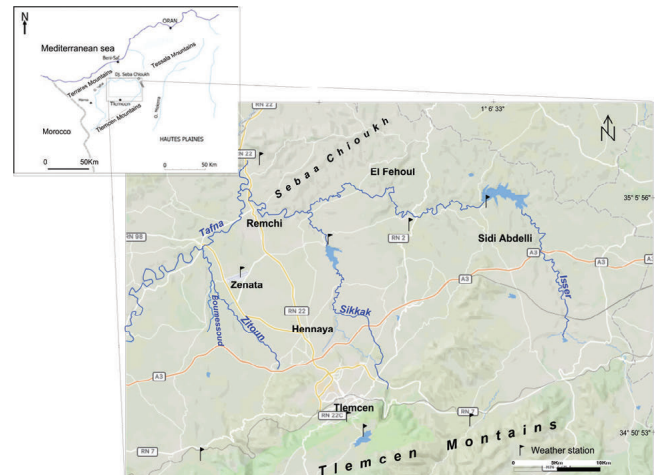


Fig. 1 - Geomorphological map of the study area (topographic base from mapcarta.com).  
Fig. 1 - Carta geomorfologica dell'area studiata (base topografica da mapcarta.com).

- Lower member: dolomitic limestone formations, 200 to 300 m thick;
- and the Upper member: dolomitic limestone formations, 120 to 150 m thick.

The second one is characterised by marl with sandstone intercalations, dating from the Serravallian, and sandstone formations dating from the Tortonian. To illustrate this, we used a lithostratigraphic detailed description of the principal units of the area (Clair, 1974; Guardia, 1975 and 1984; Benest et al., 1999; Bensaoula, 2007 and Kerzabi, 2008), together with a hydrogeological description (Fig. 3). Only the Jurassic formations can form significant groundwater reservoirs. However, several geoelectric prospectings (ENAGEO, 1979; 1983; 1987; ICOSIUM, 2002) have identified resistive layers in the Remchi plain that are associated to the Serravalian aquifer (Fig. 4), which provides a cumulative flow of 25 to 140 L/s.

### Geoelectrical data acquisition

The geological structure of North of Algeria plays a major role in the occurrence of thermal springs; the majority of thermal springs in the centre of Algeria are favoured by faults oriented NE-SW, NW-SE, EW (Nedjâi, 1987). This occurs also in Western Algeria, where the faults oriented NE-SW are connected with many thermal springs (Fekraoui, 2007). Thus, a structural overview of the Tlemcen region becomes useful to understand the thermal behaviour of groundwater.

The Tlemcen Mountains are affected by a succession of horsts and grabens oriented N50° to N70°. They are limited by normal faults, whose displacement reaches at least 500m (Collignon, 1986). There is an important system of subvertically dipping fractures, either longitudinal SW-NE to WSW-ENE (corresponding with the most frequent Tellian trend) or oblique with transverse faults N-S to NNE-SSW (Benest et al., 1999).

The recent use of geophysics (ICOSIUM, 2002) shows the existence of an important resistive formation that can be characterised by high hydraulic conductivity. This is why we carried out a geoelectric prospecting campaign. Conventional

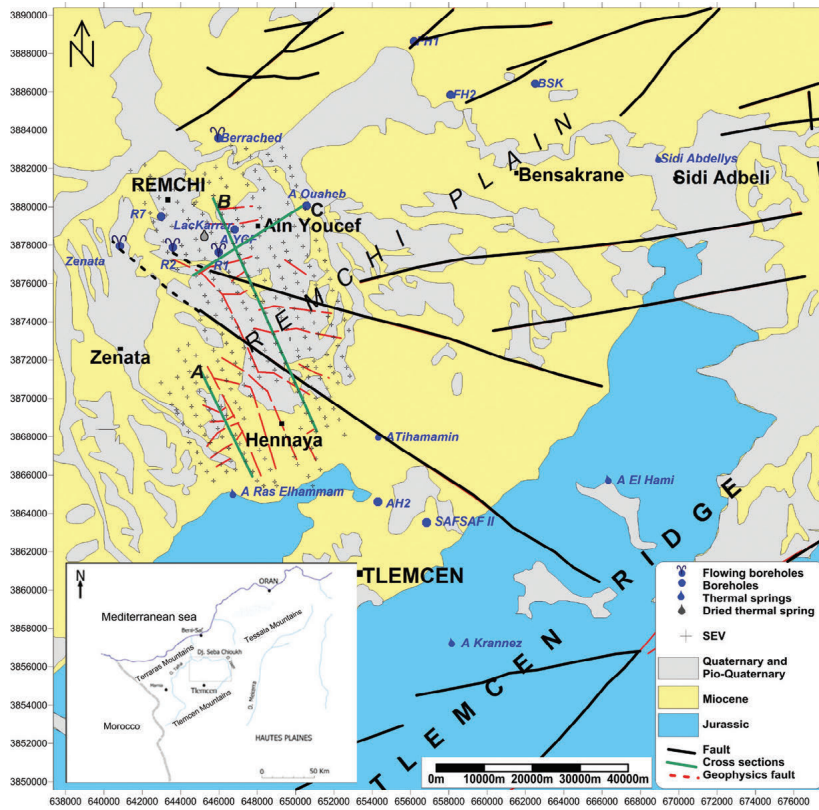


Fig. 2 - Geostructural map of the Tlemcen Mountains and the Remchi plain (ANRH, 2008, simplified).

Fig. 2 - Carta geostrutturale dei Monti Tlemcen e della Piana di Remchi (ANRH, 2008, semplificata).

Age	Geological Description	Hydrogeological description
Plio-Quaternary Upper Miocene	Alluvium, colluvial deposits recent, dune & screes Tufa	Alluvial aquifer
	Ancient limestone crust: shell limestone Stony alluvium with sand-clay cement, conglomerate Tortonian sandstone	Porous and fissured aquifer
Middle Miocene (Serravalian)	Blue grey marl with red sandstone (some cm to tens of meters)	Confined aquifer in terms of sandstone levels
Kimmeridgian	Terni dolomite Lato limestone	Karst aquifer
	Raourai marly-limestone	Karst aquifer
	Stah limestone	
	Tlemcen dolomite Zarifet limestone	
Oxfordian (upper)	Boumédiene sandstone	Aquifer moderately porous and fissured

Fig. 3 - Lithostratigraphic and hydrogeological log.

Fig. 3 - Log litostratigrafici e idrogeologici.



Fig. 4 - Outcrops of Serravalian formation in Sebba Chioukh Mountains.

Fig. 4 - Affioramento della Formazione Serravalliana nelle Montagne di Sebba Chioukh.



Schlumberger array (Fig. 5) (Telford et al., 1990) was used to perform 317 vertical electrical soundings (VES) (located in Fig. 2), with the electrode spacing (AB) being extended to a maximum of 4000 m in some of the sites. This study uses resistivity calibration values (Tab. 1), obtained both by in situ measurements near the formation outcrops, and by comparison with existing boreholes lithostratigraphic information (8 boreholes, with depth varying between 90 and 600 m). Figure 6 presents the results of modelled curve types in the following sites: Hennaya, Karrar, Zenata, Ain Youcef, Sid M'barek, RMC1, RMC2 and Berrached. This was represented by the apparent resistivity curve, the forward model and the relative 1D resistivity model.

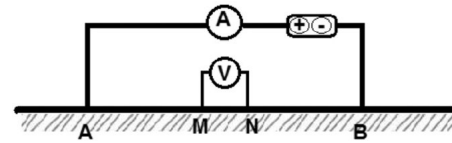


Fig. 5 - Schlumberger electrode configuration.  
Fig. 5 - Configurazione degli elettrodi Schlumberger.

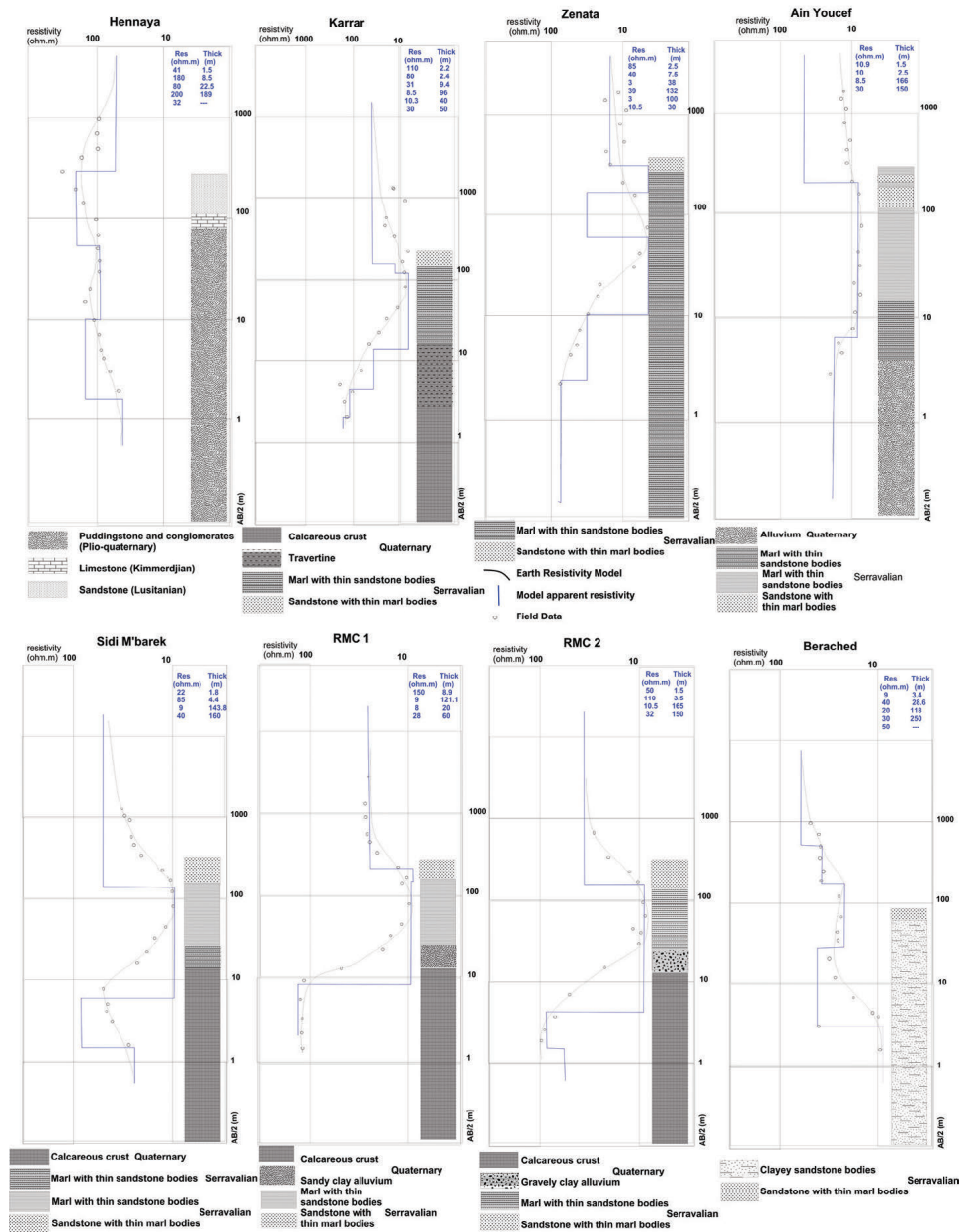


Fig. 6 - Calibration of VES with boreholes lithostratigraphic column (ICOSIUM, 2002).  
Fig. 6 - Calibrazione di VES con i logs stratigrafici (ICOSIUM, 2002).

Tab. 1 - Calibration values of resistivity (ICOSIUM 2002).

Tab. 1 - Valori di calibrazione della resistività (ICOSIUM 2002).

Formations	Resistivity (Ohm-m)	Age
Alluvium	10-280	Quaternary
Calcareous crust	23-210	
Travertine	31-480	
Pudding stone, conglomerate and sandstone	42-150	Plio-Quaternary
Yellow clay with sandstone bodies	10	
Blue marl	2-9	Miocene (Serravalian)
Blue marl with thin sandstone bodies	9-16	
Sandstone with thin marl bodies	19-77	
Limestone et dolomitic formations	90-700	Jurassic (Kimmeridgian)

**Hydrochemical data acquisition**

As part of a monitoring of the water quality of the Remchi plain, the ANRH in collaboration with the ADE (Algérienne Des Eaux) carried out a sampling campaign in March 2015 on 11 boreholes and Sidi Abdellys spring (Fig. 7). All boreholes capture the Serravalian sandstone aquifer. Temperatures and pH were measured in situ. The sample was kept cold and stored in 1.5 litre polyethylene bottles. The chemical characteristics (Cl<sup>-</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, K<sup>+</sup>, Na<sup>+</sup>, SO<sub>4</sub><sup>2-</sup>, HCO<sub>3</sub><sup>-</sup>, and NO<sub>3</sub><sup>-</sup>) were measured in the laboratory according to the analysis methods recommended by Rodier (2009). We used only the physico-chemical analyses with ionic balance values better than ±5 %. The temperatures of the outflowing waters in the study boreholes are mapped in Figure 7.

**Results and Discussions**

**Goelectrical study**

We have processed and interpreted the geoelectric data and achieved all geoelectric profiles in order to locate any anomaly in resistivity variation. These anomalies, representing faults detected by geophysical prospection, are plotted on the geological map (Fig. 2).

The combination of geoelectric cross-sections, surface and subsurface geological data (Fig. 8) allowed us to recognize three zones:

1. in the South of the Remchi plain, Miocene sandstone formations are frequent and shallow. They are in contact with the Jurassic formations (Fig. 8- A);

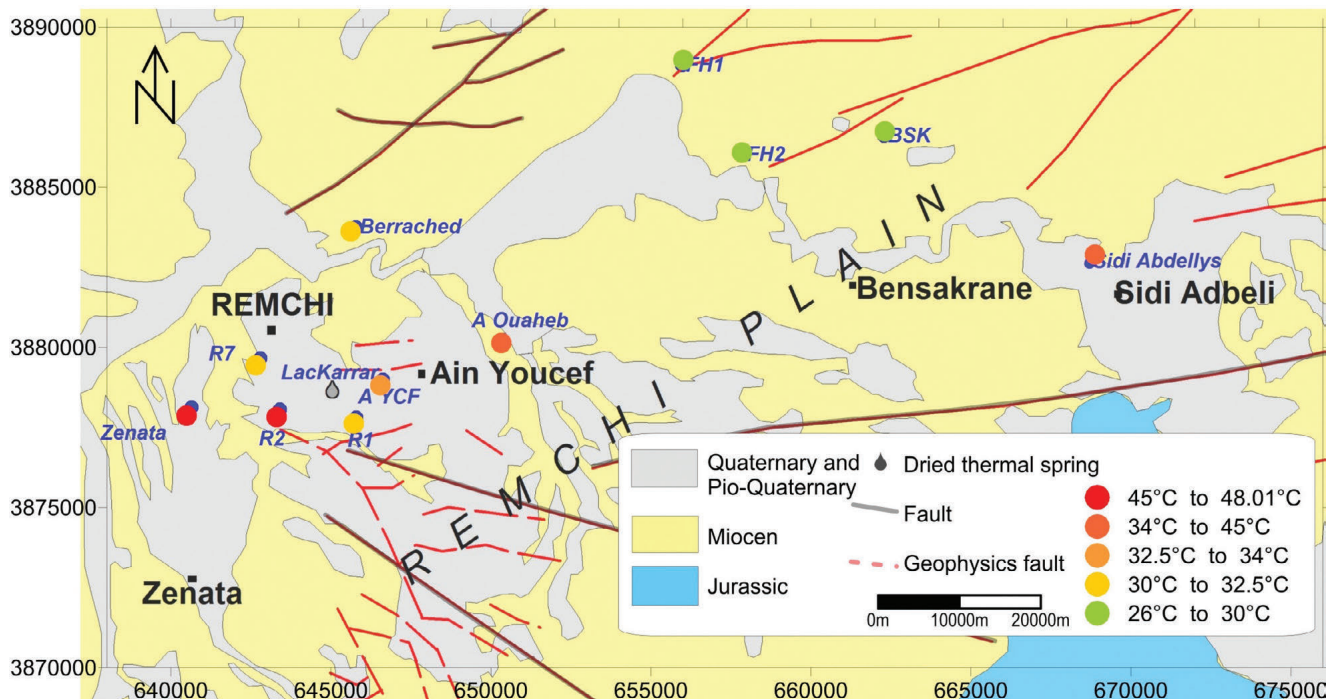


Fig. 7 - Sampling map showing the spatial water temperature variability of thermal boreholes and spring in the Remchi plain.

Fig. 7 - Carta di campionamento che mostra la variabilità spaziale della temperatura dell'acqua dei pozzi termali e delle sorgenti nella pianura di Remchi

2. in the centre of the Remchi plain, Miocene sandstone formations are absent at depths lower than 450 m. This zone constitutes a waterproof section which isolates laterally the third zone from the Tlemcen Mountains (Fig. 8- B);
3. in the Northern zone, the sandstone formations constitute thick layers with depths ranging between 50 and 300m. They constitute a deep aquifer with significant hydraulic potentiality (Fig. 8-C).

The iso-resistivity maps show the lateral and vertical variation of the apparent resistivity in the study area for five depth slices (Fig. 9). The analysis of the apparent resistivity map AB=200 m allowed us to distinguish the following areas: a) high resistive area in the South, corresponding

to Jurassic outcrops; b) intermediate resistive area in the center, occupying by the Pudding stone, conglomerate of Plio-Quaternary; c) the rest of the map is occupied by low resistivity, corresponding to Serravalian marls.

With AB spacing between 600 and 2000 m, the distribution of apparent resistivity highlights two zones (1 and 3 in Fig. 8B) that correspond well to the sandstone with thin marl layers. Resistivity values lower than 20 ohm-m characterise the center part of the map and some part of the northern west, which corresponds to the area occupied by the Serravalian conductive level (zone 2 in Fig. 8B).

The map for the last layer resistivity (AB 4000m) specifies a low resistivity in the northern part of the study area; these values might be due to the absence of sandstone layers at great depth.

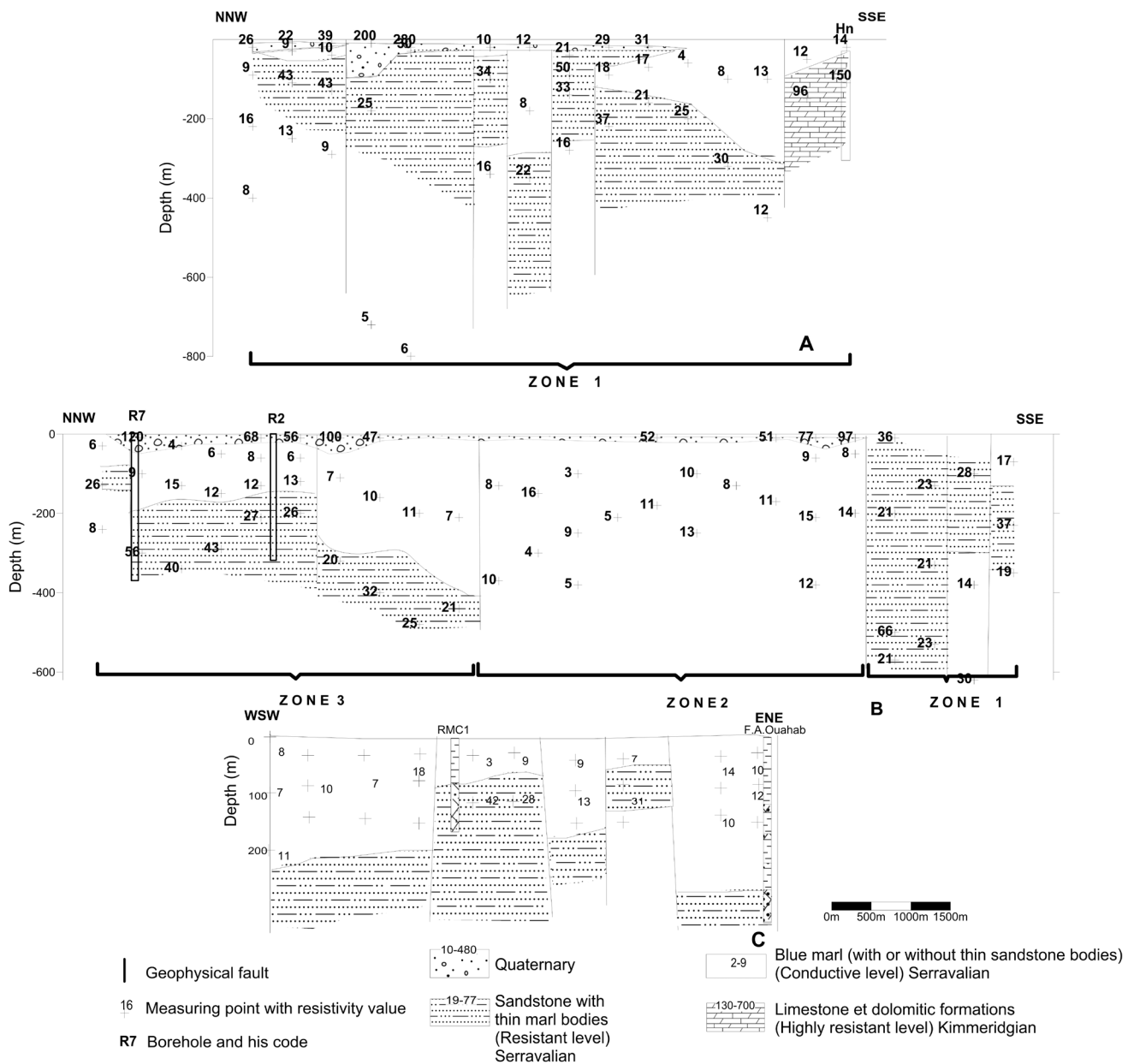


Fig. 8 - Geoelectric cross-sections.

Fig. 8 - Sezioni geoelettriche.

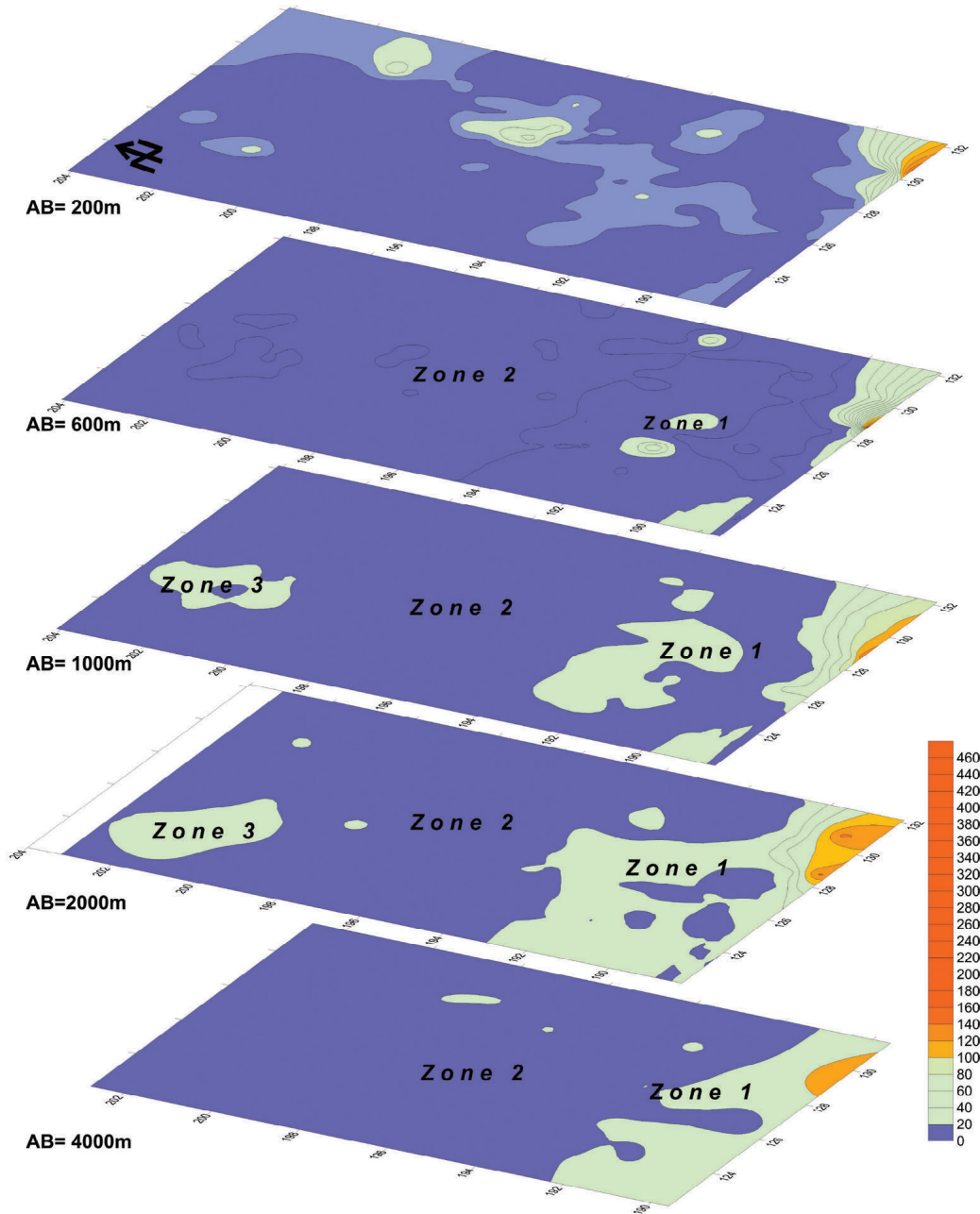


Fig. 9 - Isoresistivity maps for different AB.  
 Fig. 9 - Carte di isoresistività per diversi AB.

**Hydrochemical study**

A Piper-Kelly diagram is used to characterise the groundwater geochemistry through the identification of water type. Seven distinct grading criteria were employed, from freshwater to conservative mixing water, to characterise the seawater intrusion (Zamroni et al., 2021; Gueddari et al., 2022) or salt-water encroachment from the Sebkhia (Mhamdi et al., 2022).

Clustering Remchi plain groundwater samples in the Kelly diagram (Fig. 10) shows that Cl-SO<sub>4</sub>/Na-Ca is the dominant water facies. In general, the abundance of anions displays a Cl (40-80%) > SO<sub>4</sub> (20-60%) > HCO<sub>3</sub> (10-50%) evolution. While the abundance of cations displays a Na<sup>+</sup>K (47-85%) >

Ca (15-55%) > Mg (8-40%) evolution.

The knowledge of the chemical characteristics was used to investigate the origin of salinity in the aquifer (Dazy et al., 1987; Bianchetti et al., 1992; Abdesselam et al., 2000; Khaska et al., 2013 and Saberinasr et al., 2019). The groundwater samples from the Serravalian aquifer reveal the chloride type of groundwater; this is probably the result of exchanges with Triassic formations, which constitute the first source of chloride ions. However, the latter are unknown as outcrops in the Tlemcen Mountains (Collignon, 1986). The acquisition of sulphate could be accounted for the dissolution of gypsum which is present in the lower marls of the Miocene. It also indicates that the thermal waters have a deep flow and a



calcareous recharge area origin ( $\text{HCO}_3^-$  concentration). The content of  $^{18}\text{O}$  at Lake Karar shows that the recharge area is quite high, probably the Tlemcen mountains (Collignon, 1986). S. Abdellys spring is the only one displaying the bicarbonate type. It suggests that this aquifer is made of Jurassic carbonate rocks.

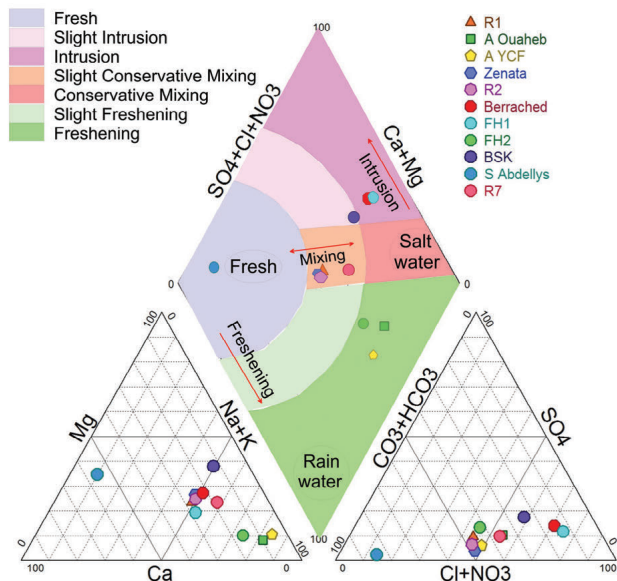


Fig. 10 - Piper-Kelly diagram of the Remchi plain water samples.  
 Fig. 10 - Diagramma Piper-Kelly dei campioni di acqua della pianura di Remchi.

Durov diagram has been used to identify hydrochemical processes in groundwater, classifying different types of water mixing, the presence of ions, and reverse ion exchange processes. The geochemical processes occurring in the groundwater of the study area were further verified by Durov's plot (Fig. 11). It showed that the groundwater samples are plotted in fields 4 and 5. The fitting of samples in field 4 showed the occurrence of  $\text{Ca}^{2+}$  and  $\text{SO}_4^{2-}$  as the dominant ions in the water, and they are mainly due to the gypsum dissolution process. Some samples are plotted in field 5, which indicates that water exhibits simple dissolution or mixing occurring between two or more different facies (no dominant cation or anion) (Lloyd & Heathcoat, 1985). The recent study in the Hennaya plain (Plio-quadernary aquifer) (Yebdri, 2023)

Tab. 2 - Pearson correlation matrix for Serravalian aquifer parameters in the Remchi plain.

Tab. 2 - Matrice di correlazione di Pearson per i parametri dell'acquifero serravaliano nella pianura di Remchi.

Variables	Ca	Mg	K	Na	Cl	SO <sub>4</sub>	HCO <sub>3</sub>	NO <sub>3</sub>
Ca	1							
Mg	0.672	1						
K	-0.226	0.411	1					
Na	0.498	0.505	0.061	1				
Cl	0.842	0.781	0.063	0.856	1			
SO <sub>4</sub>	0.689	0.866	0.312	0.819	0.914	1		
HCO <sub>3</sub>	-0.295	-0.134	0.051	0.446	0.038	0.236	1	
NO <sub>3</sub>	0.541	0.380	0.033	0.190	0.384	0.518	0.080	1

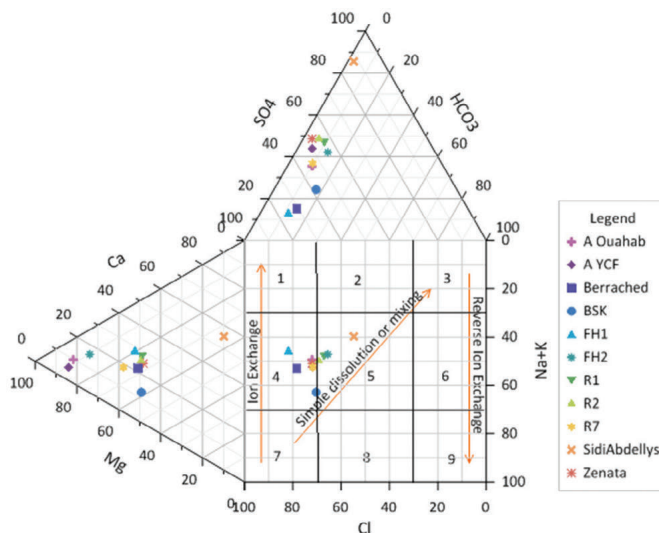


Fig. 11 - Durov plot illustrating hydrochemical processes involved in groundwater in study area.  
 Fig. 11 - Grafico di Durov che illustra i processi idrochimici coinvolti nelle acque sotterranee nell'area di studio.

showed that 98% of samples belong to group 5. These waters are considered as mixed water, which confirms that they are mixed with a recent freshwater recharge, highlighting the contribution of the Northern recharge area (Sebâa Chioukh Mountains) to the aquifer recharge.

**Statistical analysis**

**Correlation analysis**

Bivariate correlation technique is initially applied as the first assessment of the hydrochemical origin of the facies. Admitting that for 10 degrees of freedom, a correlation factor  $R^2 > 0.815$  implies a statistically significant degree of correlation, Table 2 displays a good correlation between  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Na}^+$  and  $\text{Cl}^-$ ,  $\text{SO}_4^{2-}$ , suggesting that most elements have an evaporite origin.

**Principal Component Analysis (PCA)**

Principal Component Analysis is largely used to analyse hydrochemical data of karst aquifers (Bakalowicz, 1979; Blavoux & Mudry, 1986; Razack & Dazy, 1990 and Hamit,



2012), and it is considered an efficient statistical tool to infer groundwater origin (Mudry 1987 and Marín Celestino 2018). It is applied to groundwater quality data to extract principal factors corresponding to the different sources of hydrochemistry variation (Helena et al., 2000; Mouser et al., 2005 and Belkhiri et al., 2010), and contamination sources (natural and anthropogenic; De Andrade et al., 2008), especially in semi-arid zones (Yousfi, 2014). PCA of the hydrochemical data will clearly define the element markers of the Miocene sandstone aquifer in the confined zone and the flow conditions of groundwater (Fig. 12). Eight hydrochemical variables were considered:  $Mg^{2+}$ ,  $Ca^{2+}$ ,  $K^+$ ,  $Na^+$ ,  $SO_4^{2-}$ ,  $Cl^-$ ,  $HCO_3^-$  and  $NO_3^-$ . Two factors (F1 and F2) explaining together 72% of total variance allowed us to classify the variables into two distinct groups (Fig. 12): the first cluster is defined by  $Mg^{2+}$ ,  $Ca^{2+}$ ,  $Na^+$ ,  $SO_4^{2-}$ ,  $Cl^-$ . Explaining 54% of the variance. These tracers explain the dissolution of Triassic formation in deep and evaporite rocks that are present in the basis of Miocene.  $HCO_3^-$  make alone the second group, with 18.13% of variance indicating a carbonate origin. This factor plan is typical of an evaporate-carbonate typical behaviour. Two tracers are poorly represented in the factor plan 1x2.  $K^+$ , which is generally typical of clayey media, defines alone F3, and  $NO_3^-$ , typical of anthropogenic origins, defines alone F4. The data set is highly dominated by halite and gypsum, two major minerals of evaporate contents.

The analysis of the projection for S. Abdelllys spring and boreholes in the factor plane F1x2 (Fig. 12) shows that the measurement points can be subdivided into three groups:

- The first group: characterised by strong mineralization in boreholes FH1 and Berrached (north of Remchi plain). This mineralization is connected to the lithology of the aquifer;

- The second group presents high rates of bicarbonate corresponds to the boreholes FH2, BSK, A. Ouahab and A. YCF. These boreholes are located along a fault, favouring upward circulation of deep bicarbonate waters.
- The third group: These are the least mineralized waters, compared to the other groups, Zenata, R7, R2, R1 and S. Abdelllys.

**Geothermal conditions**

To estimate the depth of the deep aquifer, we calculated the thermal anomaly of the water according to Collignon (1986), by correcting the air and water temperatures. The results are presented in table 3, with a minimal depth reached by water ranging between 277 and 867 m. FH1, FH2 and BSK are characterised the lower values, where the sandstone formation is shallower. In the West (Zenata and R2), water temperature is 45 and 48°C, respectively, which can be due to a deeper aquifer and a rapid water rise favoured by faults. The rising waters at 31°C feeding Lake Karar (currently dry) showed a similar context of emergence (Fig. 13), favoured by a fault. This can be explained by the three boreholes, heavily mineralized with bicarbonates (A. Ouahab, FH2, A. YCF and BSK) which are aligned and located near a geological fault which should cross Lake Karar. Around this spring, we can clearly distinguish a travertine massif (3 km<sup>2</sup>) with a few tens of meters of thickness. At the core of this massif, mesothermal water emerges. This basin existed until the beginning of the 20th century, but its age can be dated from the Middle Pleistocene.

The relationship between thermal anomaly and the mean screen depth (Fig. 14) presents a linear correlation, displaying a geothermal gradient of 4.8°C/100m, that agrees with

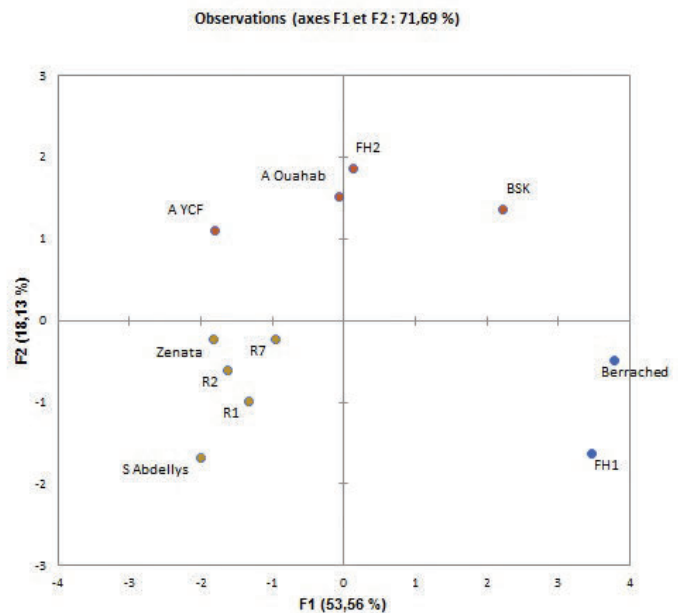
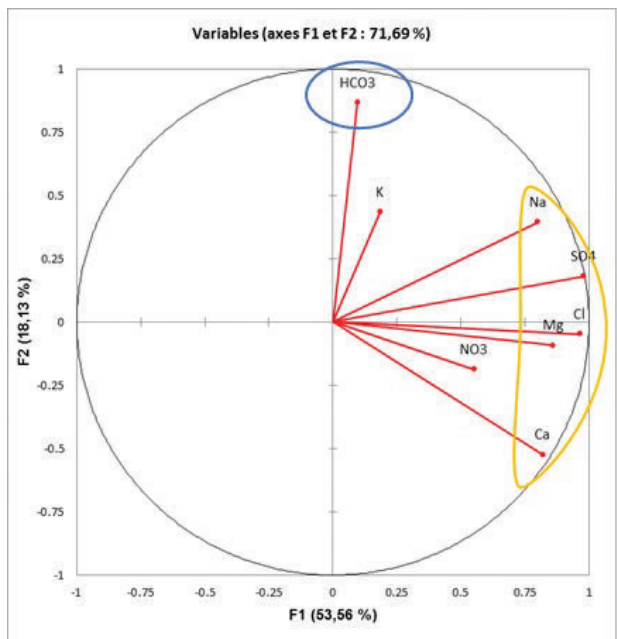


Fig. 12 - Correlation between chemical parameters and factors F.  
 Fig. 12 - Correlazione tra parametri chimici e fattori F.

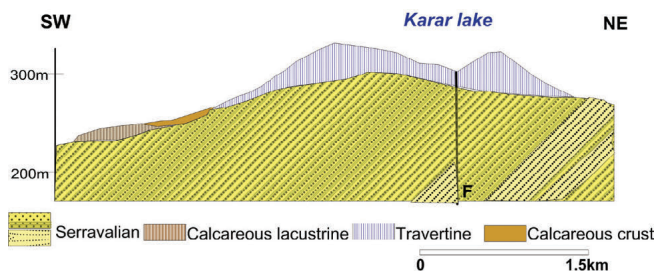


Fig. 13 - Geological cross section of Karar Lake .

Fig. 13 - Sezione geologica del Lago Karar.

the calculated one in the Oran meseta region (Blavoux & Collignon, 1986).

Tab. 3 - Boreholes characteristics.

Tab. 3 - Caratteristiche dei pozzi.

Location	Alt (asl)	Correction of altitude (Z*0.55/100)	Tair= 17.29- Calt	Water Temperature Tw (asl)	Mean screen depth (P)	Expected temperature (asl)	Thermal anomaly Tw-Tair	Estimated depth of the deep aquifer
	m	Calt	°C	°C	m	Tair+P*0.037	°C	m
R2	248	1.4	15.9	48	241	24.8	23.2	867
Zenata	219	1.2	16.1	45	318	27.8	17.2	781
A Ouahab	203	1.1	16.2	37	240	25.1	11.9	563
S. Abdellys	350	1.9	15.4	34	0	15.4	18.6	504
A YCF	250	1.4	15.9	32.5	214	23.8	8.7	448
R1	275	1.5	15.8	32	149	21.3	10.7	438
R7	210	1.2	16.1	32	303	27.3	4.7	429
Lac Karar	327	1.8	15.5	31	0	15.5	15.5	419
Berrached	150	0.8	16.5	30	154	22.2	7.8	366
BSK	480	2.6	14.7	26	170	20.9	5.1	307
FH2	300	1.7	15.6	26	149	21.1	4.9	280
FH1	280	1.5	15.8	26	87	19.0	7.0	277

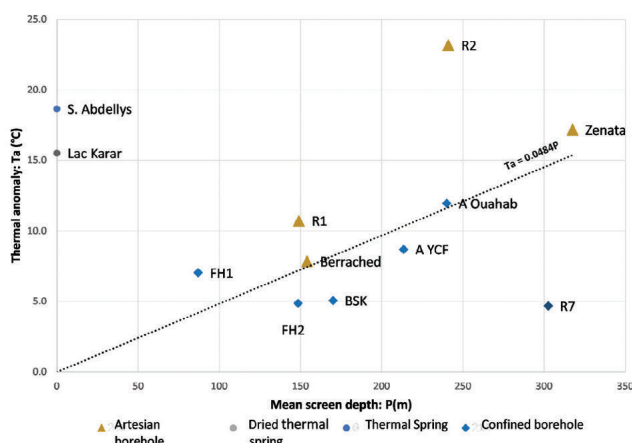


Fig. 14 - Thermal anomaly vs. mean screen depth.

Fig. 14 - Anomalia termale vs. profondità media dei filtri.

### Hydrogeological model

Geoelectric cross-sections, in combination with surface geological data and hydrochemical analyses, allowed us to build a conceptual hydrogeological model of groundwater flow from recharge zones of the Tlemcen Mountains to the thermal aquifer (Fig. 15).

This cross-section was realised by the projection of outcropping formations in the Tlemcen sheet of the geological maps (1/50,000), boreholes, lithostratigraphic column, and geophysics information (faults and location of sandstone formations), in order to complete the knowledge about geological setting and the regional groundwater flow. We identify three zones: the first, where the sandstone formations are in contact with Jurassic formations of the Tlemcen Mountains, which take part in lateral water inflow. The second zone is characterised by the absence of sandstone formations at more than 450 m depth. The third zone represents the confined thermal sandstone aquifer. The structure of this

zone plays an important role for vertical water inflow in the West of the Remchi Plain (Zenata, R7, R2, etc.); there is an increasing depth of the sandstone aquifer in the East, where the water temperature is low (FH1, FH2, Berrached).

Physical-chemical analyses help us to understand the hydrodynamic behaviour of groundwater. The thermal water in Remchi plain spends a long time in the sandstone formations, explained by a substantial vertical path in the South, from the Jurassic recharge area to the aquifer, deeper than 450 m. The hypothesis of halite dissolution from deeper Triassic formation could explain the origin of chloride concentrations in the Remchi plain groundwater. The limestone outcrops of Jurassic formations are the principal recharge area of the thermal aquifer, which provides bicarbonates in all water samples. In addition, the proportional relationship between water temperature and depth of the aquifer top corroborates the above-mentioned results.

The flow map of the Tlemcen Mountains (Fig. 16) shows

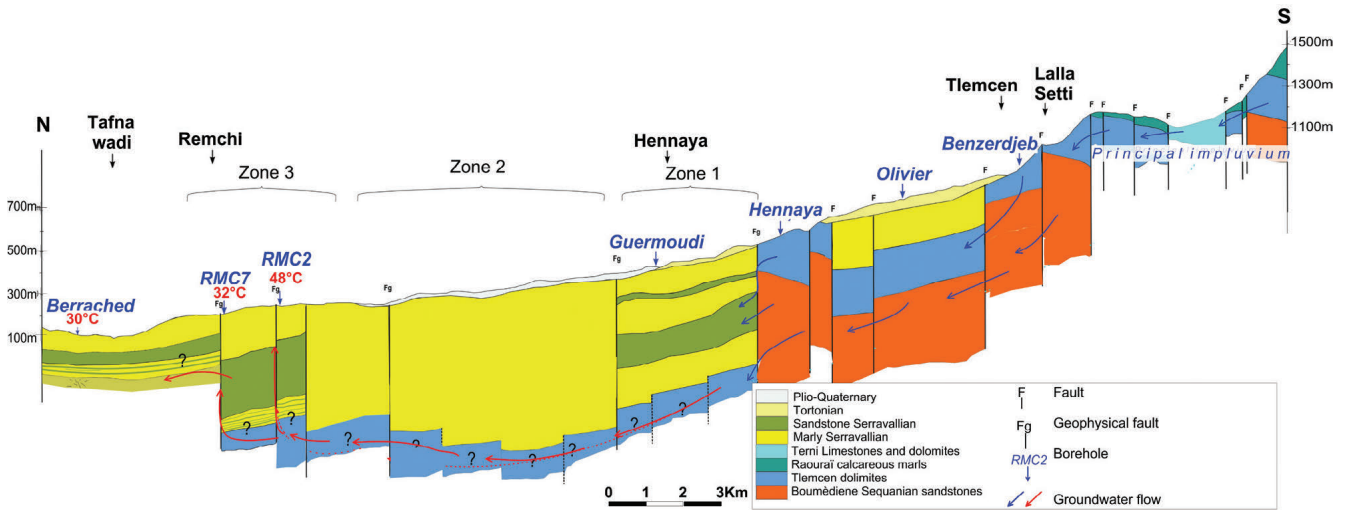


Fig. 15 - Hydrogeological model of groundwater.  
 Fig. 15 - Modello idrogeologico delle acque sotterranee.

that the main direction of flows in the Northern foothills is towards units buried under the Miocene formation. The aquifer system extends under low permeable or impermeable cover; there are piezometric indicators to specify the flow direction. This map can explain artesian flow in Remchi Plain.

**Conclusions**

Multi-criteria analyses of data about the deeper aquifer in Remchi plain, show the important relationship between water circulation and physical-chemical parameters; in this study case, the high temperature of water is explained by a significant path to the aquifer, deeper than 450 m. The

interpretation of the different geological sections as well as all the apparent iso-resistivity maps allowed us to identify the geometry of Serravalian sandstone-marl layers. We have a better understanding of vertical and lateral hydraulic connections between sandstone or marly sandstone aquifers and Jurassic dolomitic limestone formations.

Geospatial analysis of hydrochemical parameters indicated spatial variation and heterogeneities in water quality analyses. In the study area, heat water ascent is caused by normal faults crossing the plain to the SW and NE. This water has a deep path, contributing to its mineralization. The Miocene marls can be the origin of high concentrations of gypsum elements, whereas the halite elements indicate that sulphates are likely

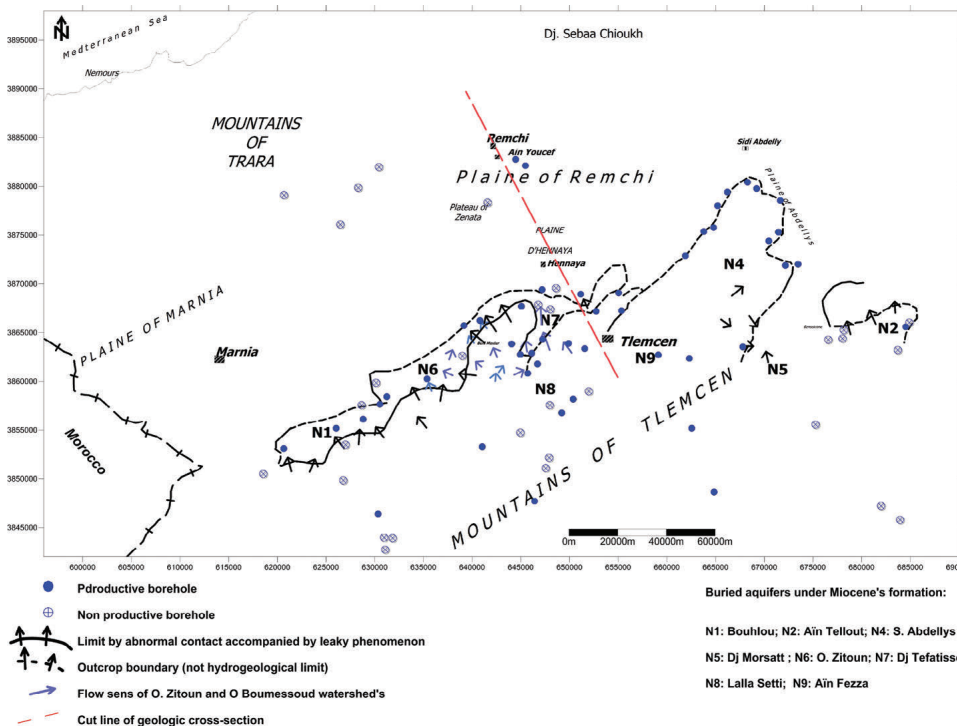


Fig. 16 - Flows of karst aquifers (Northern foothills of Tlemcen mountains) (ANRH, 1995 modified).

Fig. 16 - Flusso degli acquiferi carsici (versante settentrionale ai piedi dei monti Tlemcen) (ANRH, modificato nel 1995).

to originate from dissolution within Triassic formations. Some studies have been conducted on the groundwater quality of 50 water points in the northern foothills of the Tlemcen Mountains: O. Zitoun and O. Boumessoud (Touati & Nebia, 2019), Sabra and Béni Senous (Boukacem & Derbal, 2019) and Tlemcen Urban Group (Amar Belarbi & Benslimane, 2023) show that nitrate concentrations range from 12 to 98 mg/L (Beni Mester<sup>2</sup>). We deduce that the shallow location of these waters favours contamination by fertilizers and/or discharges of urban wastewater, while the low nitrate content in the Remchi plain waters (less than 10 mg/L) can be explained by a deeper origin.

It should also be noted that the thermal waters of the Zouia deep Liassic aquifer (Western foothills of the Tlemcen Mountains) have Sodium bicarbonate, Sodium chloride and Sodium sulphate facies (Bensaoula et al., 2005; Bensaoula & Adjim, 2006; Benkhelifa & Lamouri, 2018). The origin of chlorides and sodium may be the leaching from Triassic formations (Bensaoula & Adjim, 2006).

The results yielded by this survey show the necessity of preparing a more detailed study (isotopic analyses), which should help in exploiting this thermal aquifer.

The hydrogeological model built in this study, based on geological, hydrogeological and geophysical information, identified the extension of the thermal aquifer and explained the groundwater circulation from the Jurassic recharge zone to the sandstone aquifer.

#### Competing interest

All authors, declare no competing interests.

#### Author contributions

Collection of data and interpretation of results, Rachid Kerzabi and Abelhamid Achachi; data processing, interpretation of results and writing-original draft preparation, Somia Yousfi; visualization and supervision, Jacques Mudry, Mustapha Bensalah and Bernard Collignon.

All authors have read and agreed to the final version of the manuscript.

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